

Assimilative Model for Ionospheric Dynamics Employing Delay, Doppler, and Direction of Arrival Measurements from Multiple HF Channels

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The Ionospheric Reconstruction Problem: Tikhonov Method

 $N(\mathbf{r},t) = N_0(\mathbf{r},t)Q(u(\mathbf{r},t)); \quad \text{i.e. } Q(u(\mathbf{r},t)) = e^{u(\mathbf{r},t)}$ $U = \{u(\mathbf{r},t)\}$ $Y \approx M[U]$

Y is the set of measured values obtained via various ionospheric measurements (such as TEC data, HF oblique propagation delay)



-The pseudo-covariance *P* matrix is defined in such a way that the stabilizing functional tends to take on larger values for unreasonably behaving solutions ("reasonable" \Leftrightarrow "smooth").

-The nonlinear optimization problem is solved iteratively (Newton-Kontorovich).



HF Oblique Propagation Data within GPSII

-Simulated values of measured data can be obtained for any ionospheric model U via numerical ray tracing (RT).

-This defines the non-linear functional of measurements M[U]

Ray Tracing Equations

Hamiltonian Formulation of RT Equations [Haselgrove, 1957, Jones, 1975]





HF Oblique Propagation Data within GPSII

-The non-linear inverse problem is solved iteratively as a sequence of linear problems. At the iteration *n* the non-linear functional M[U] is approximated by a linear operator *L* as follows

$$M[U] = M[U_{n-1}] + L(U - U_{n-1}) + o(||U - U_{n-1}||) \qquad \Leftrightarrow L = \delta M / \delta U$$

- L is the Ray Path Response (RPR) operator

-The Ray Path Response operator L is estimated using the extended RT equations – the equations augmented with the linearized ray-tracing equations

Extended RT Equations





Evaluation of the Ray Path Response Operator L





Evaluation of RPR within the 3-D ionospheric inversion problem can be performed with little computational burden because it is reduced to computation of several one-dimensional integrals

For HF data the main computational burden remains associated with the classical ray homing task. This task precedes computation of RPR

Numerical representation of the RPR is a sparse matrix with nonzero elements occupying only nodes of the spatial grid that are adjacent to the ray trajectories that connect receiver and transmitter. Matrix operations with RPR are not a substantial computational burden as we take advantage of the sparse character of RPR.



Test with Range-Doppler Data Set



- Florida Collection (August 13, 2013)
 Range/Doppler Data
 - Receiver at Vero Beach, FL
 - Multiple transmitter sites
 - Three hour collection of 3 KHz bandwidth FMCW waveform



NWRA Range-Doppler Data Compared to GPSII Fit

Since 1984





NWRA GPSII Solution with Range-Doppler Data





Assimilation of Oblique lonograms (along with TEC and VI data)

Geography of DSTO data sources employed in this test*





Comparison of GPSII results for 2 links

GPSII runs with oblique ionogram reproduce observed OI details (yellow) for both assimilated (left) and test (right) propagation links

Both links shown at 01:45. All synthetic OIs are extraordinary-ray traces.



Impact of OI data on GPSII Model

Vertical cut through the model at latitude -18 degrees



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- The theoretical framework for incorporating HF channel probe data (propagation delay, angles of arrival, Doppler shift) into ionospheric inversion algorithms has been developed
- Capabilities to assimilate data from HF channel probes and oblique ionograms have been added to GPSII
- Performance and validation of the algorithm are addressed in the companion paper by L.J. Nickish



Backup

Evaluation of the Ray Path Response Operator L

GPSII Solution with Range-Doppler Data and the GPS TEC data

Reconstruction of a TID from Simulated OTHR Data

$$\delta U\Big|_{t} = \delta U(t_{m}) \frac{t_{m} - t}{t_{m} - t_{m-1}} + \delta U(t_{m-1}) \frac{t - t_{m-1}}{t_{m} - t_{m-1}}; \quad \delta \dot{U}\Big|_{t} = \delta U(t_{m}) \frac{1}{t_{m} - t_{m-1}} - \delta U(t_{m-1}) \frac{1}{t_{m} - t_{m-1}}$$