Direct Forcing of the Thermosphere-lonosphere by Small-Scale Gravity Waves of Lower Atmospheric Origin

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Overview

Science: Atmospheric Coupling, Waves, & General Circulation Models

2 Gravity Wave Effects on the Upper Atmosphere

3 Possible effects on the ionosphere

4 Summary & Conclusions

5 Open science questions









Vertical coupling: Meteorological effects

UPPER ATMOSPHERE (THERMOSPHERE-IONOSPHERE)

(COMMUNICATION, PLANETARY MISSIONS,

SATELLITE ORBITS, SPACE TRAVEL)

Significant wave-flow interactions

MIDDLE ATMOSPHERE	Wave-flow interactions
Lower Atmosphere	Weather/Wave sources → Spectrum of waves

Vertical coupling: Meteorological effects

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Significant wave-flow interactions





Gravity waves



Noctilucent clouds - Gravity waves

What are internal gravity waves?

Gravity waves (GWs)

- Buoyancy oscillations
- Primary generation in the lower atmosphere
- Propagate upward, interacting continuously with the atmospheric flow, e.g., via nonlinearity and viscosity

Small-scale GWs

- Important for the general circulation of the whole atmosphere system
- Challenge: unresolved & parameterized in general circulation models (GCMs)
- Small-scale \Leftrightarrow large-scale

Problem: Previous parameterizations have ignored the effects of GWs in the upper atmosphere.



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Dynamical equations

$$\frac{\partial \mathbf{u}}{\partial t} = \overline{\left(-\frac{1}{\rho}\nabla p - (\mathbf{u}\cdot\nabla)\mathbf{u} + \nabla(\nu\nabla\cdot\mathbf{u}) - 2\boldsymbol{\Omega}\times\mathbf{u} + \mathbf{g} - \nu_{ni}(\mathbf{u}-\mathbf{v}_i)\right)}$$
(1)



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Dynamical equations

$$\frac{\mathbf{a}_{res}}{\partial t} = -\frac{1}{\rho} \nabla p - (\mathbf{u} \cdot \nabla) \mathbf{u} + \nabla (\nu \nabla \cdot \mathbf{u}) - 2 \Omega \times \mathbf{u} + \mathbf{g} - \nu_{ni} (\mathbf{u} - \mathbf{v}_i) \quad (1) + \mathbf{a}_{gw}$$

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Dynamical equations

Prof. Yiğit

Ero

$$\frac{\partial \mathbf{u}}{\partial t} = \underbrace{-\frac{1}{\rho} \nabla \rho - (\mathbf{u} \cdot \nabla) \mathbf{u} + \nabla (\nu \nabla \cdot \mathbf{u}) - 2 \Omega \times \mathbf{u} + \mathbf{g} - \nu_{ni} (\mathbf{u} - \mathbf{v}_i)}_{+ \mathbf{a}_{gw}} (1)$$

+ \mathbf{a}_{gw}
 $\frac{\partial T}{\partial t} = Q_{res}$

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$$+ \mathbf{a}_{gw}$$
$$\frac{\partial T}{\partial t} = \Omega_{res} + \Omega_{gw}$$
(2)

Gravity Wave Effects in the Atmosphere – Year 2008 AERONOMIST proviting wover are generated at high-latitudes and propagate equatoriard. Zever struggheric small-scale Give are negligible, as they are small perturbation," 90 km GWs get dissipated anyways before those reach the thermosphere 15km METEOROLOGIST Who is right?

A whole atmosphere parameterization

GW whole atmosphere effects

So, how do we get a physics-based representation of the GW contributions a_{gw} and Q_{gw} to the energy and momentum balance in the whole atmosphere?



A whole atmosphere parameterization

Click Here Full Article JOURNAL OF GEOPHYSICAL RESEARCH, VOL. 113, D19106, doi:10.1029/2008JD010135, 2008

Parameterization of the effects of vertically propagating gravity waves for thermosphere general circulation models: Sensitivity study

Erdal Yiğit,¹ Alan D. Aylward,¹ and Alexander S. Medvedev²

Received 16 March 2008; revised 8 July 2008; accepted 22 July 2008; published 8 October 2008.

[1] A parameterization of gravity wave (GW) drag, suitable for implementation into general circulation models (GCMs) extending into the thermosphere is presented.

- "Extended parameterization" (Yiğit and Medvedev 2013)
- Small-scale (subgrid-scale) GWs in GCMs (Yiğit et al. 2008)

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- Input: Initial gravity wave activity at a given source level (e.g., ~15 km)

Gravity wave propagation & dissipation

$$\overline{u'w_i'}(z) = \overline{u'w_i'}(z_0)\rho(z_0)\rho^{-1}(z)\tau_i(z)$$

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(4)

Output

GW induced dynamical (\mathbf{a}_{gw}) and thermal effects (Q_{gw})

Modeling framework: GCM + paramaterization

GCM

A General Circulation Model that extends into the thermosphere (Coupled Middle Atmosphere Thermosphere Model-2)



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GW representation

A GW parameterization that accounts for GW thermospheric dissipation (Spectral nonlinear GW parameterization (Yiğit et al. 2008))

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GW effects in the thermosphere-ionosphere

#



Simulated mean GW effects with the Coupled Middle Atmosphere Thermosphere Model-2 ((CMAT2), Yiğit 2009) implementing the Yiğit et al. 2008 GW parameterization (Yiğit et al. 2009, Figure 3)

Gravity wave effects in Earth's upper atmosphere



Prof. Yiğit

Global Scale Modeling

Gravity wave effects in Earth's upper atmosphere



GW heating/cooling in Earth's upper atmosphere a) 75°S, 140°W 240 3000 230 2000 cooling 220 1000 700 210 500 (km 200 200 190 100 650 180 50 Altitude heating 170 10 650 160 0[K/d] 150 -10 600 140 -50 550 130 -100 500 50 400 450 400 120 -200 350 350 300 650 110 -500 200 200 100 -1000 03Z 06Z 09Z 12Z 15Z 18Z 21Z 00Z 00Z -2000 **22JUN** 23JU] -3000 1985 UT

Simulated GW heating/cooling (color) and temperatures (contour) (Yiğit and Medvedev 2009, Figure 3)



et al. 2014, Fig 3

GW propagation during SSWs

Science question

What is the role of gravity waves in the vertical coupling between the lower and upper atmosphere during sudden stratospheric warmings?

A minor SSWs



GW effects during SSWs



GW propagation into the thermosphere during the different phases of the minor warming (Yiğit et al. 2014, Figure 2abc)

lonospheric effects



International Beacon Symp 21 / 25

Ionospheric effects

Prof. Yiğit



- GW dynamical effects of up to
 - \pm 120 m s⁻¹ day⁻¹ in the lower thermosphere
 - 200 to 400 m s⁻¹ day⁻¹ in the high-latitude thermosphere

are much stronger than the middle atmospheric effect of GWs, which is around 40 m $s^{-1} \ day^{-1}.$

Thermospheric GW effects
 comparable to ion drag

Science goal

What are the effects of GWs on the upper atmosphere?



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- During sudden warmings, thermospheric GW effects increase significantly.

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Conclusions

- GCMs \rightarrow physical insight into coupling processes (waves, storms)
- Application of a GCM implementing the Yiğit et al. 2008 scheme: powerful tool
- Gravity effects should be taken into account in the energy and momentum budget studies of the thermosphere-ionosphere system

COSPAR

COSPAR C2.2 Wave Coupling and Consequences in the Whole Atmosphere

• Moscow, 2014, MSO \rightarrow NEXT: Istanbul, 2016, MSO

SCOSTEP's VarSITI Program

 Working Group Leader: Coupling by Dynamics, ROSMIC (Role Of the Sun and the Middle atmosphere/thermosphere/ionosphere In Climate) program

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Future: Some Open Science Questions

- What are the influences of meteorological processes in shaping the state, evolution, and variability of planetary upper atmospheres?
- What is the impact of space weather on atmospheres and vertical coupling processes?
- What is the significance of <u>wave-wave interactions</u> for atmospheric vertical coupling?

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ADDITIONAL MATERIAL

General circulation modeling



- Goal: to predict the future state of the atmospheric circulation from knowledge of its present state by using numerical approximations techniques to the dynamical equations
- Solve coupled conservation equations in a global grid
- Differential equations \rightarrow algebraic difference equations
- From the initial state, integrate the equations in time to get future state.