## Three-Dimensional Modeling of High-Latitude Scintillation Observations

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# Summary

- Scintillations observed in high-rate GPS signals provide a means of studying ionospheric behavior beyond the resolution of EISCAT incoherent scatter radar. At GPS frequencies (L1: 1575 MHz, L2: 1228 MHz), intermediate-scale irregularities (approximately 0.1-10 km) are responsible for diffractive scattering.
- Case study: the ionospheric electron density profile is observed by EISCAT incoherent scatter radar (Tromsø) along the same line-ofsight as a scintillating GPS signal
- EISCAT large-scale (10s of km) densities constrain a 3D irregularity model with a multiple phase screen propagation algorithm.
- The observed signal is modeled and likely characteristics of the underlying ionospheric irregularities are estimated

## **Observational Approach**

- Experiment at Tromsø (66.73 N, 102.18 E): EISCAT UHF antenna aimed at GPS satellite PRN 23, beam tracking every 5 minutes. Ground scintillation monitor takes 50 Hz observations
- Scintillations observed just after 20:00 UT (close to midnight magnetic local time) on 17 October 2013



#### **Observations**



- E-region density enhancement occurs in the auroral zone around midnight MLT (20:05 UT).
- Spike in the detrended 50 Hz GPS carrier phase at the same time.
- Line-of-sight velocities drop to ~100 m/s
- Large (1500 K) electron temperature enhancement characteristic of electron precipitation
- No corresponding ion temperature enhancement

## **Observations**

- Density enhancement associated with a phase spike but no observable intensity
- 50 Hz data detrended using 3<sup>rd</sup>-order polynomial and 6<sup>th</sup>-order Butterworth 0.1-Hz high pass filter.
- Power spectral densities calculated using Welch's method with a Hamming window, eight segments and a 50% overlap



# Modeling propagation

- Satellite-beacon lonospheric scintillation Global Model of the upper Atmosphere (SIGMA) is adapted here
- Deshpande et al. [2014] developed the model and performed a comprehensive parametric sensitivity study
- Multiple phase screens constructed to represent signal scattering caused by ionospheric irregularities
- Geometry modification: Z along line-of-sight, effective drift velocity is along Y



# Modeling irregularities

Field-aligned irregularity spectrum P(k) is based on the formulation of *Costa and Kelley* [1977]:

 $P(k) = a \gamma \sin(3\pi/\gamma) / (4\pi^2 k_0^3) \Delta N^2 \cdot \{ (1 + (k_x^2 + k_y^2 + a^2 k_z^2) / k_0^2)^{-\gamma/2} \}^{-1}$ 

 $k = (k_{x'}, k_{y'}, k_{z'})$ : spatial wave number vector,  $\gamma$ : spectral index, a: axial ratio,  $\Delta N$ : root-mean-square density fluctuation,  $k_0$ : outer scale wavenumber, z': magnetic field direction.

- Grid defined between 95-175 km altitude (110 200 km range)
- Mean electron density N specified to match EISCAT, but fractional fluctuation density \(\Delta N/N\) allowed to vary. Effective velocity set to 300 m/s, X/Y extent is 3 km.
- We found a > 1 leads to intensity scintillation enhanced above what is observed, so we set a = 1 and therefore γ = 4.2 (matching observed spectral value)
- Spectral irregularities do not reproduce large (3 radians p-2-p) spike, so kilometer-scale irregularity added at 20:05:25
- 50 random realizations account for 'chance' variability

# Model irregularity configuration



## **Model-Observation comparison**



## Model-observation comparison





- △N/N is clearly not a constant in this case. Standard deviation varies between 5 - 25 % of mean value specified by EISCAT
- Absence of intensity scintillation observed on line-of-sight approximately 25° off B. This is explained as a consequence of high spectral index (4.2) and low axial ratio (1:1)
- The new geometry dramatically reduces computation times to approx. real time and reproduces the observations



#### References

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